

Mesoscale and microscale modeling of the Martian atmosphere

A. Spiga

and many co-authors



CPS Kobe mini-workshop
February 5, 2013

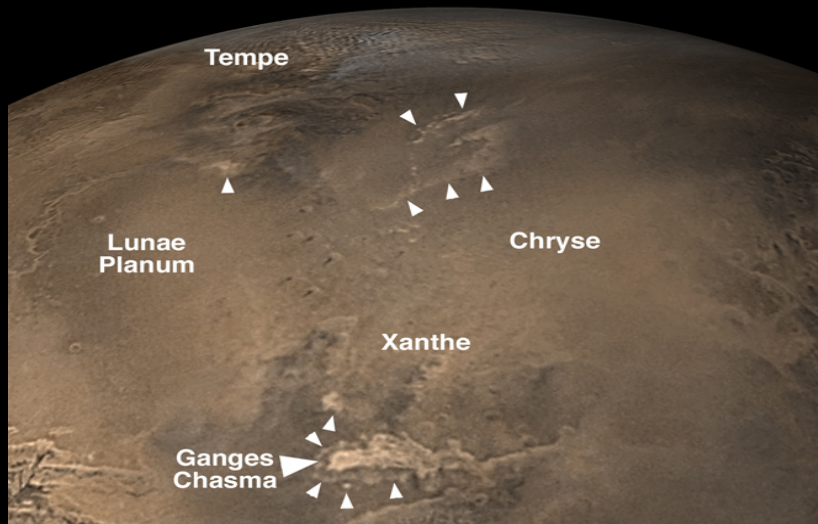
Outline

- 1 Methodology
- 2 Slope winds
 - Tharsis volcanoes
 - Polar regions
- 3 Boundary layer convection
 - Vortices
 - Radiative control
- 4 Gravity waves
 - Mesosphere
 - Troposphere
- 5 Rocket dust storms
- 6 Conclusion

Outline

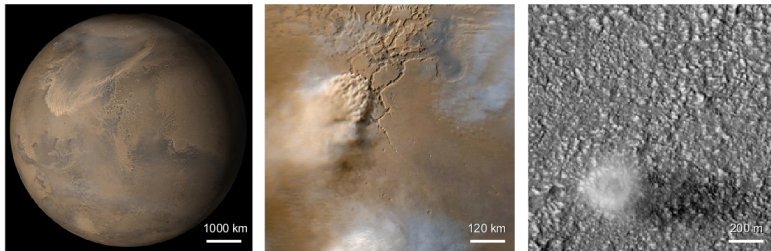
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The Martian mesoscale “zoo”



[MGS/MOC imagery, Malin Space Science Systems, 02/2002]

Echelles spatiales et modèles adaptés



... Dust fronts ... Regional dust storms ... Local gusts ... Dust devils ...

10000 km 1000 km 100 km 10 km 1 km 100 m 10m 1m

Global Circulation Models

Mesoscale Models

Large-Eddy Simulations

[Spiga and Lewis, 2010]

Mars Global Circulation Model: LMD-MGCM

LMDz dynamical core

integration of conservation laws for momentum, mass, energy, tracers

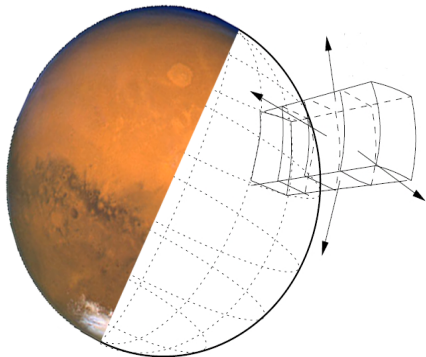
LMD Mars physics

radiative transfer (dust and CO₂), soil model, vertical mixing, microphysics (H₂O and CO₂), lifting/sedimentation, chemistry

MGS dataset

topography, thermal inertia, albedo dust scenario

Grid spacing ~ 200 km



[Forget et al., JGR 1999]

Mars Mesoscale Model: LMD-MMM

WRF dynamical core

integration of conservation laws for momentum, mass, energy, tracers

LMD Mars physics

radiative transfer (dust and CO₂), soil model, vertical mixing, microphysics (H₂O and CO₂), lifting/sedimentation, chemistry

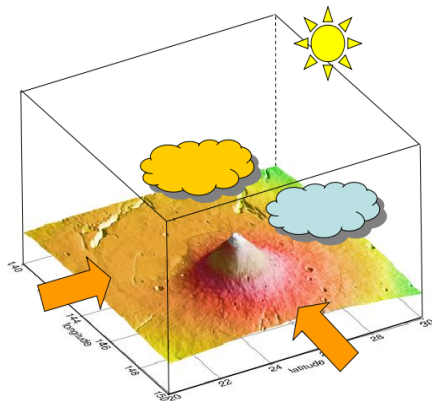
LMD Mars GCM fields

initial and boundary conditions

MGS hi-res dataset

topography, thermal inertia, albedo dust scenario

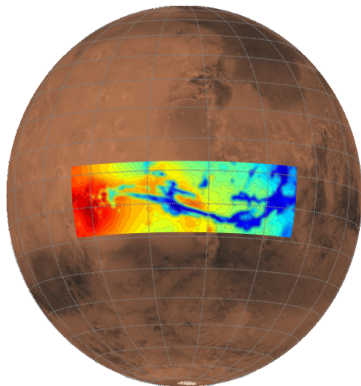
Grid spacing $\sim 10 - 1$ km



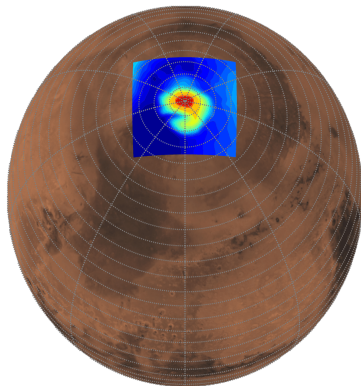
[Spiga and Forget, JGR 2009]

Mesoscale domains

Valles Marineris



North polar regions



Mars Large-Eddy Simulations: LMD-LES

WRF dynamical core

integration of conservation laws for momentum, mass, energy, tracers

LMD Mars physics

radiative transfer (dust and CO₂), soil model, vertical mixing, microphysics (H₂O and CO₂), lifting/sedimentation, chemistry

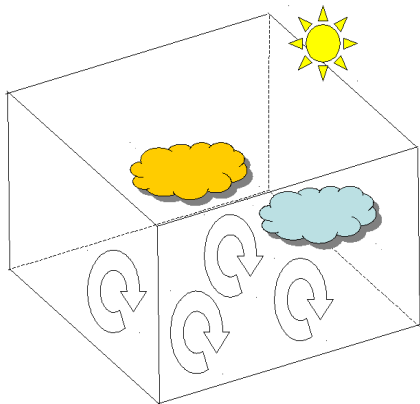
LMD Mars GCM fields

initial profiles only, periodic boundaries

MGS hi-res dataset

topography, thermal inertia, albedo prescribed dust scenario

Grid spacing $\sim 100 - 10$ m



[Spiga et al., QJRMS 2010]

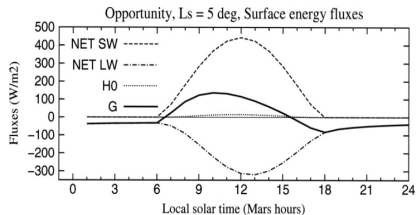
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Energy budget for Martian surface

Surface energy budget

$$F_{\text{LW}} + F_{\text{SW}} = G + H_s + LE$$



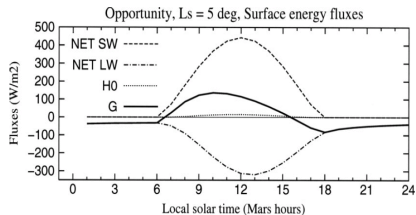
[Savijärvi and Kauhanen, QJRMS 2008]

Energy budget for Martian surface

Surface energy budget

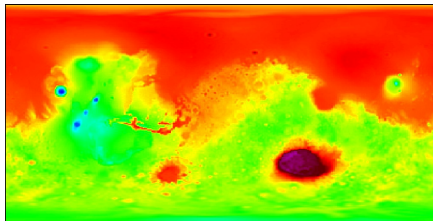
$$F_{\text{LW}} + F_{\text{SW}} = G + H_s + LE$$

→ radiative equilibrium

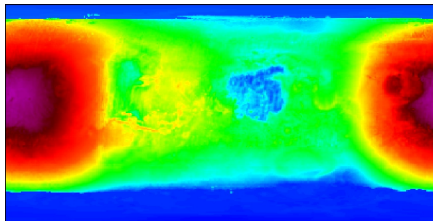


[Savijärvi and Kauhanen, QJRMS 2008]

Topography



Surface temperature



[outputs from the UK Mars GCM]

Thermal structure around a mountain on Earth



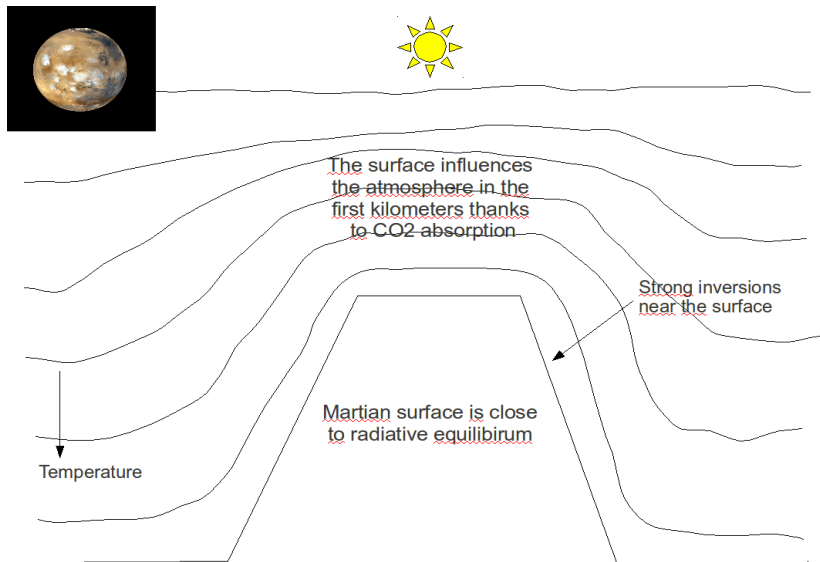
The surface influences
the atmosphere only hundred
meters above the surface

Moderate inversions
near the surface

Temperature

Terrestrial surface is in
equilibrium with the
atmosphere : sensible
heat flux, latent heat flux

Thermal structure around a mountain on Mars



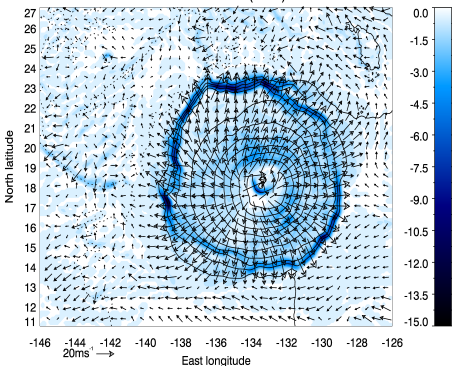
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Katabatic and anabatic winds

Nighttime downslope

Winds 10m ABG (m s^{-1})

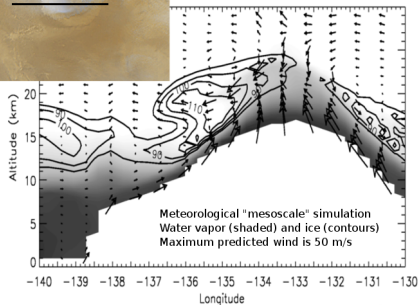


LT = 02:00am / $L_s = 173^\circ$ / $dx = 6\text{km}$ [single] / Uniform $T_1 = 85 \text{ J m}^{-2} \text{ s}^{-0.5} \text{ K}^{-1}$

Daytime upslope

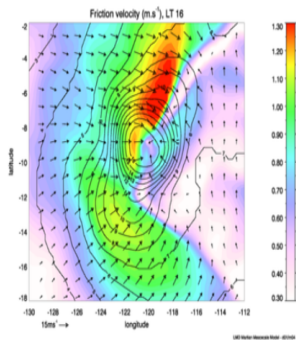
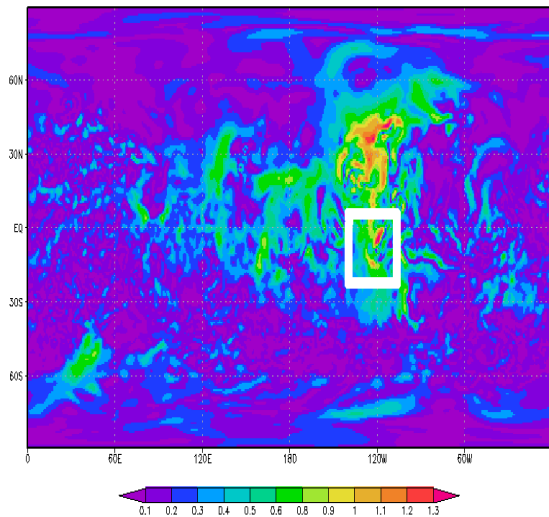


Olympus Mons Summer cloud



[Spiga and Forget JGR 2009; Spiga et al. Icarus 2011]

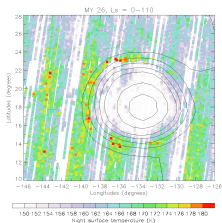
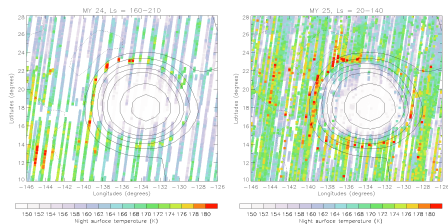
Anabatic winds: hi-res GCM vs. mesoscale



[Spiga and Lewis 2010]

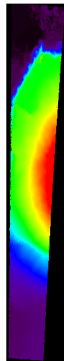
Surface temperature at night, Olympus Mons

MGs / TES

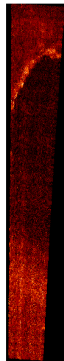


OMEGA

Olympus #8933
SEA: -65° local time: 22h30 ls: 190



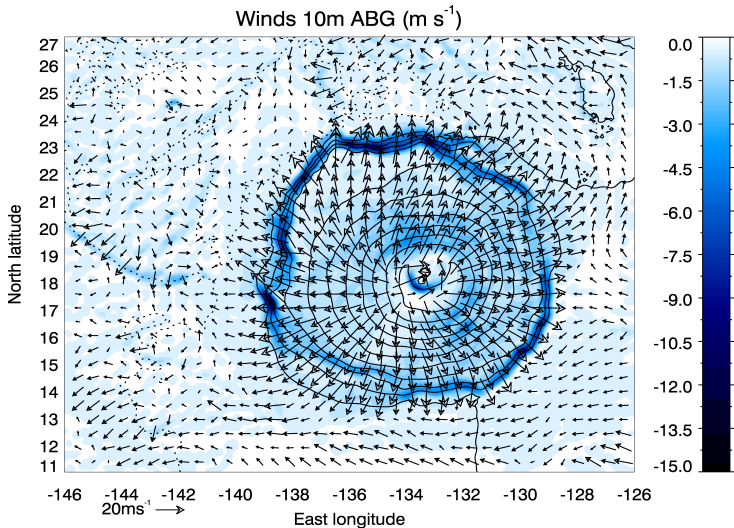
Altimetrie Mola



Omega 5 μm

[Spiga et al., Icarus 2011; Gondet and Langevin, pers. comm.]

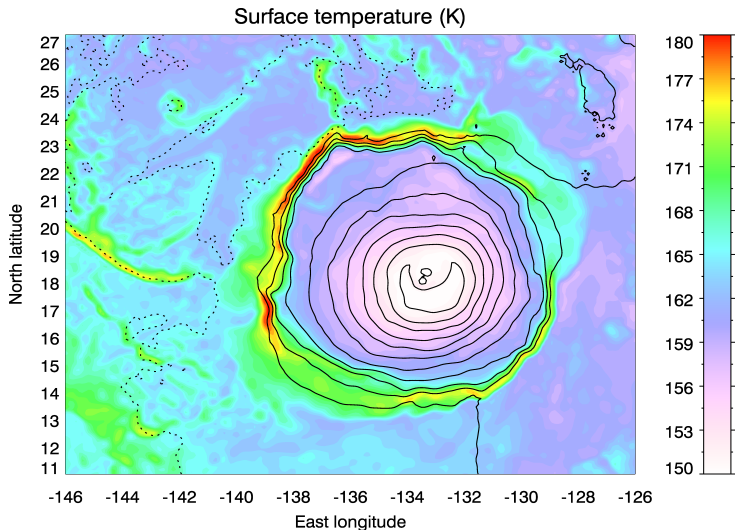
Katabatic winds in Olympus Mons and Lycus Sulci



LT = 02:00am / Ls = 173° / dx = 6km [single] / Uniform TI = $85 \text{ J m}^{-2} \text{ s}^{-0.5} \text{ K}^{-1}$

[Spiga and Forget, JGR 2009; Spiga et al., Icarus 2011]

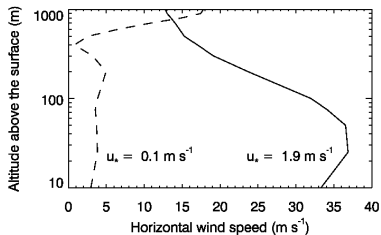
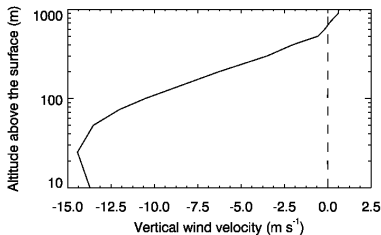
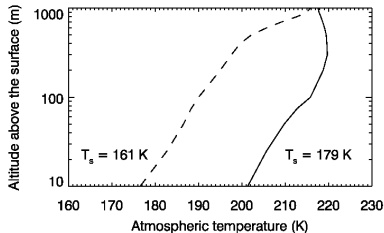
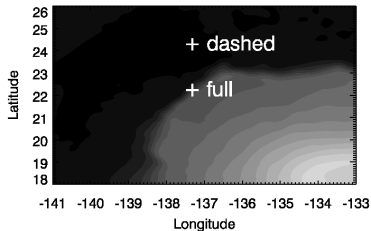
Nighttime “warm katabatic ring”



LT = 02:00am / Ls = 173° / dx = 6km [single] / Uniform TI = 85 J m⁻² s^{-0.5} K⁻¹

[Spiga et al., Icarus 2011]

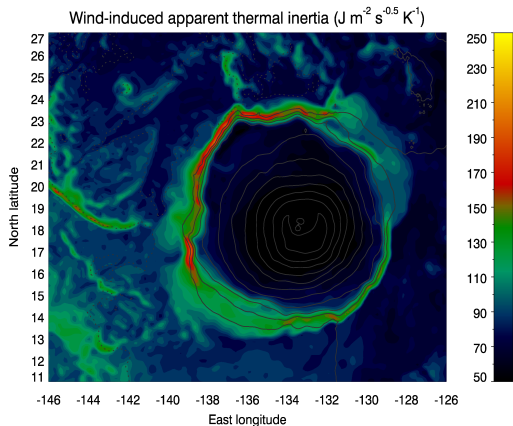
Plain vs. steep slopes: typical near-surface profiles



[Spiga et al., Icarus 2011]

Artefacts of thermal inertia

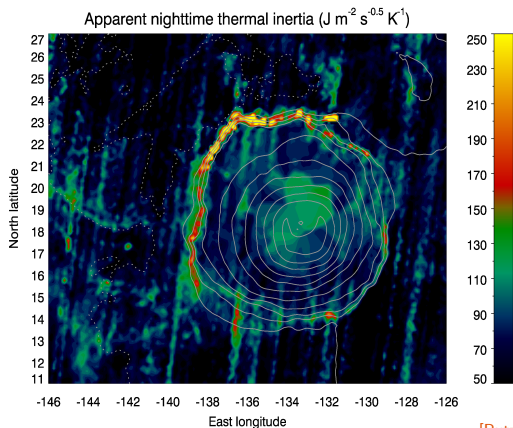
$$\epsilon \sigma T_S^4 = I \sqrt{\frac{\pi}{\tau}} \frac{\partial T_g}{\partial \zeta} \Big|_{\zeta=0} + \mathcal{F}_{\text{IR}} - H_S$$



[Spiga et al., *Icarus* 2011]

Artefacts of thermal inertia

$$\epsilon \sigma T_S^4 = I \sqrt{\frac{\pi}{\tau}} \frac{\partial T_g}{\partial \zeta} \Big|_{\zeta=0} + \mathcal{F}_{\text{IR}} - H_S$$

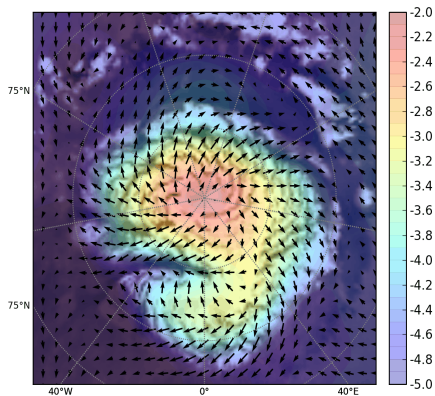
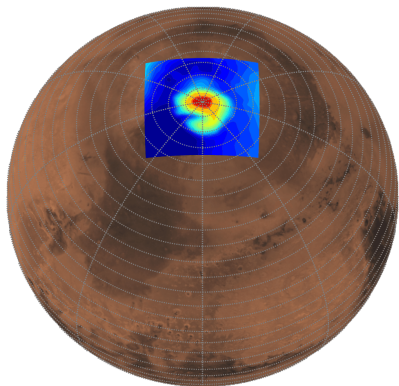


[Putzig and Mellon, Icarus 2007]

Outline

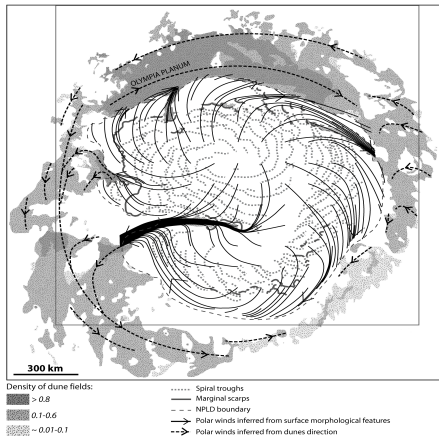
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Near-surface regional winds: northern polar cap

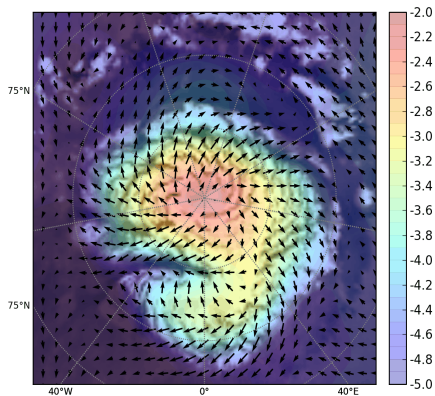


Near-surface regional winds: northern polar cap

Frost streak and dune mapping



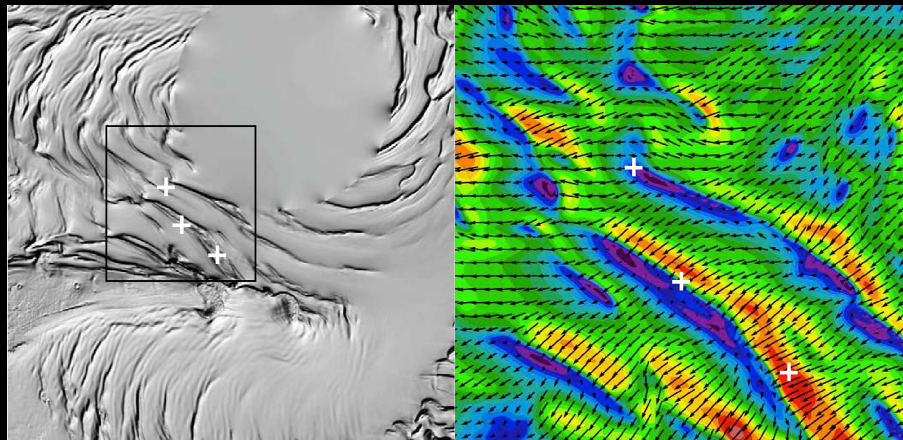
Mesoscale modeling



[Data: Howard Icarus 2000 and Massé et al. EPSL 2012]

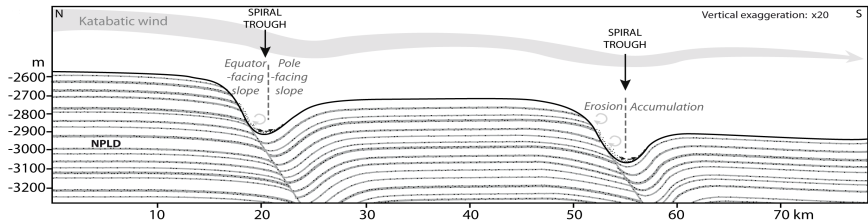
Troughs influence katabatic winds

Results from LMD mesoscale modeling with resolution 2 km

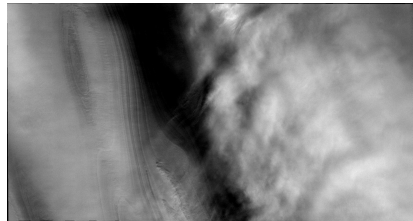
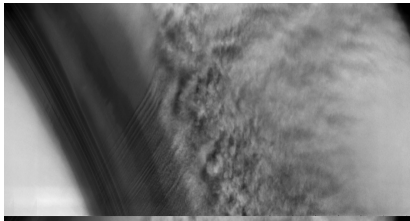


[Smith et al. submitted to JGR]

Trough migration & katabatic winds



[Massé et al., EPSL, accepted]



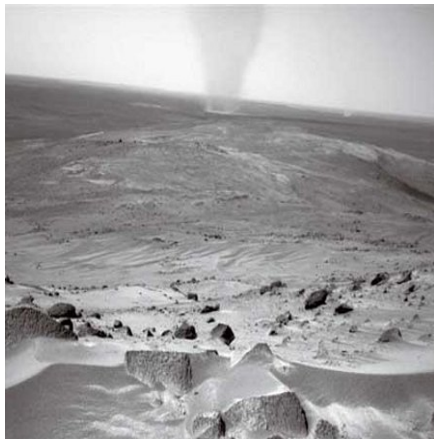
[Smith and Holt Nature 2010]

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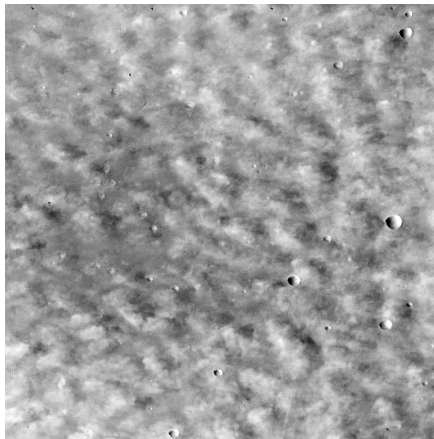
Mars PBL phenomena: Imagery

Dust devils



[Pancam on Spirit Rover]

Cloud streets

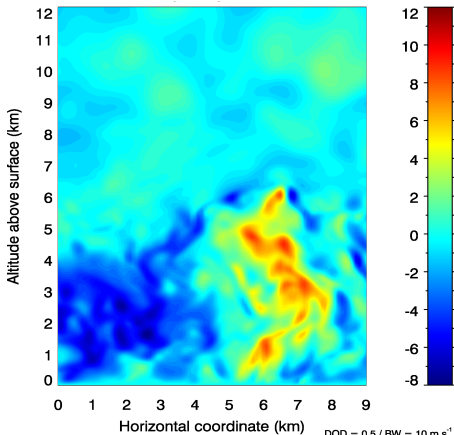


[Mars Orbital Camera on Mars Global Surveyor]

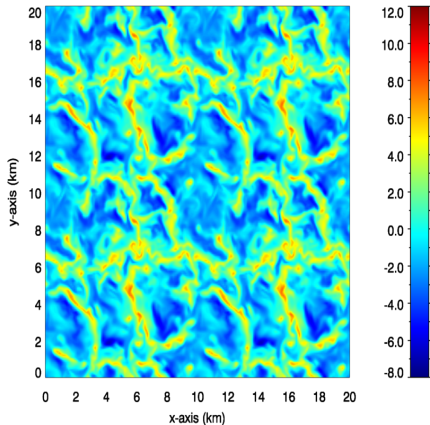
Turbulent convection in daytime boundary layer

Simulated through Large-Eddy Simulations [LES]

↓ Vertical ↑



← Horizontal →



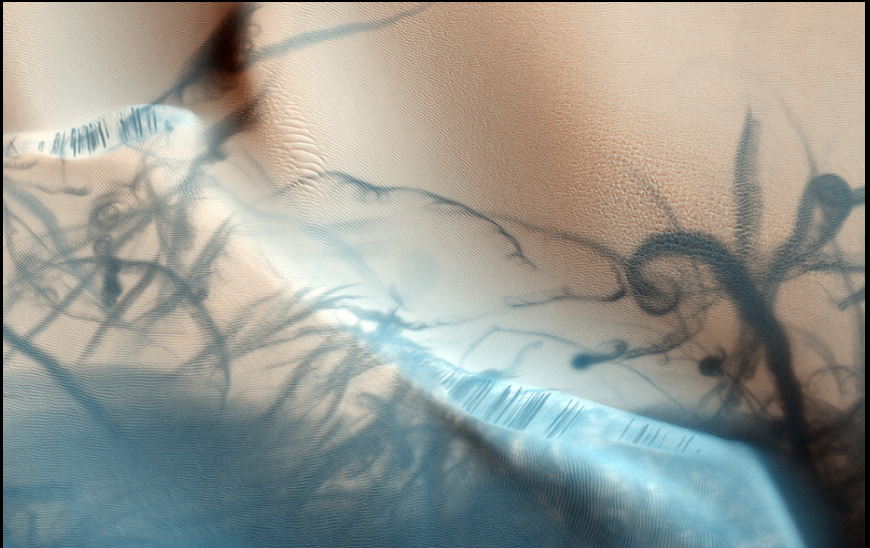
[Meridiani simulation for Exomars risk assessment]

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Dust devils observed by Spirit

Graffiti martiens ! Champ de dunes Arabia Terra

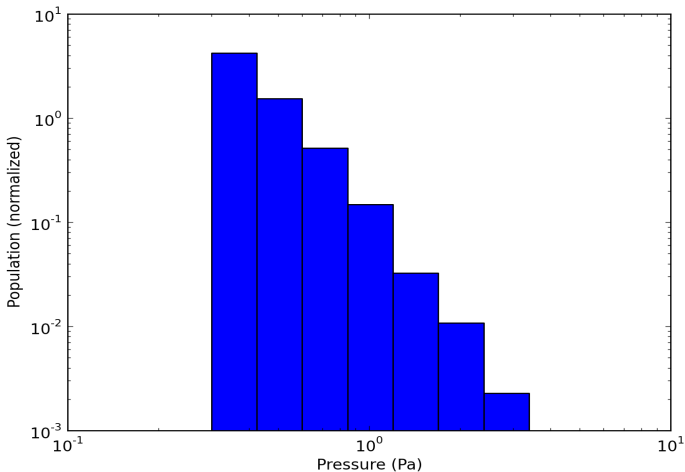


[HiRISE, Mars Reconnaissance Orbiter, 2009]

Convective vortices in LES

Distribution of pressure drop in 10 m LES

Histogram with logarithmic axes and bins

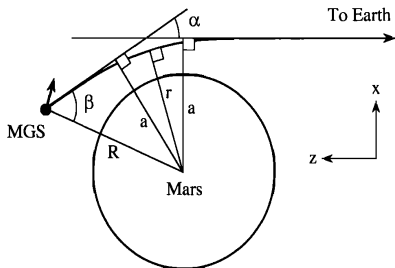
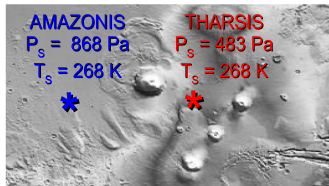


Note: nice power law fit with exponent 3.3

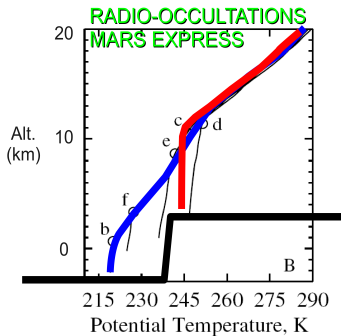
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Mars Express radio-occultations

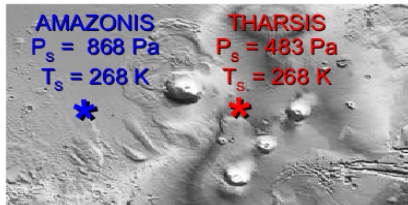


[Hinson et al., 1999]

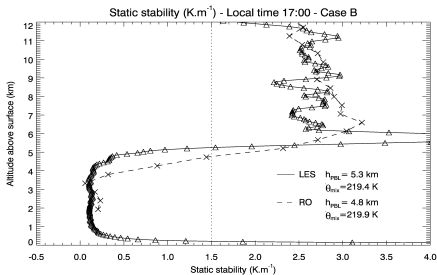


[Hinson et al., 2008]

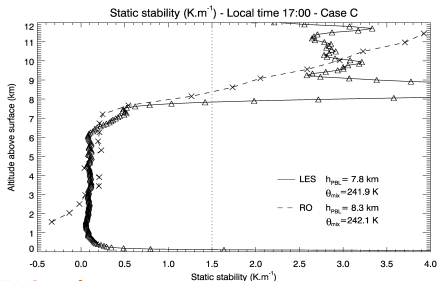
BL depth variability: observations vs. models



Lower plains [Amazonis]

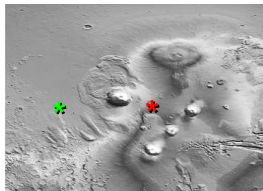


Higher plateaus [Tharsis]



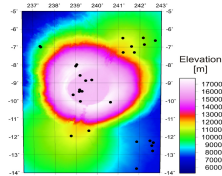
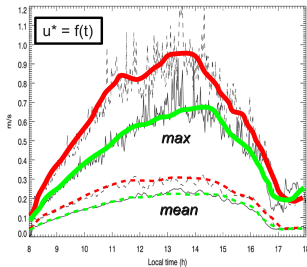
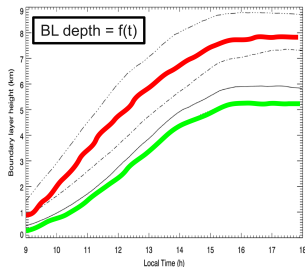
[Spiga et al., QJRMS 2010]

Existence of dust devils in high-altitude plateaus



AMAZONIS
TI = 50 tiu / A = 0.3
855 Pa

THARSIS
TI = 60 tiu / A = 0.3
483 Pa



[Reiss et al., LPSC 2009]

[Spiga and Lewis, Mars Journal 2010]

Energy budget, bottom of mixed layer

(free convection conditions)

$$c_p \frac{\partial \theta}{\partial t} = \left(\frac{p_0}{p} \right)^{R/c_p} \left[\mathcal{J}_{LH} + \mathcal{J}_{LW} + \mathcal{J}_{SW} \right] - c_p \frac{\partial \langle w' \theta' \rangle}{\partial z}$$

Mars

$$\frac{\partial \theta}{\partial t} \sim \left(\frac{p_0}{p} \right)^{R/c_p} \frac{\mathcal{J}_{LW}}{c_p}$$

Earth (arid terrains)

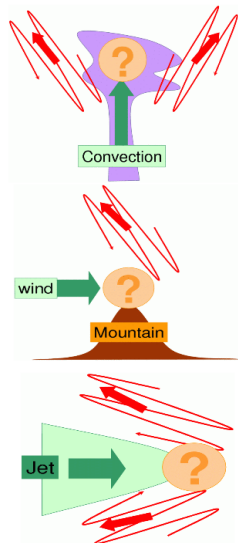
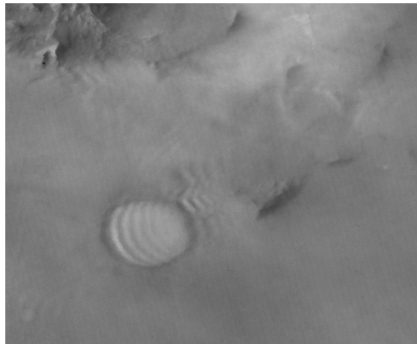
$$\frac{\partial \theta}{\partial t} \sim - \frac{\partial \langle w' \theta' \rangle}{\partial z}$$

[Spiga et al., QJRMS 2010]

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Gravity Waves [GWs] and their sources



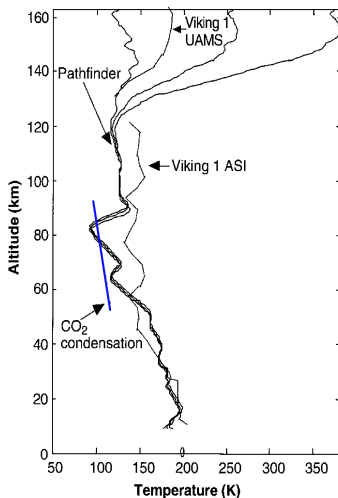
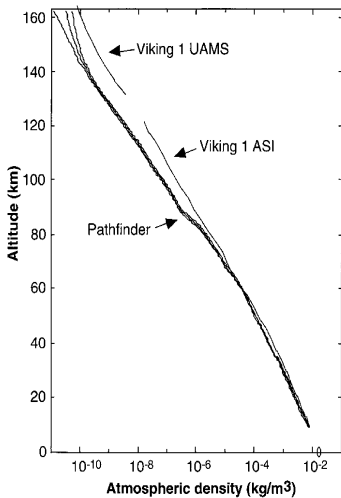
[MGS/MOC Image]

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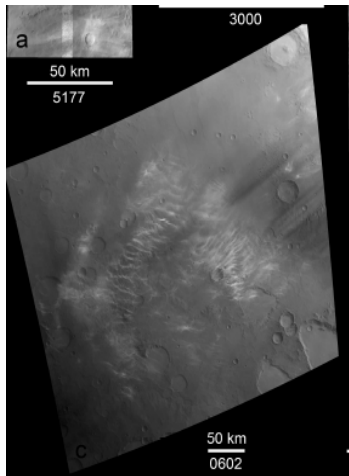
GW events observed in entry profiles

Viking [Seiff and Kirk 1977] Pathfinder [below] MGS and ODY [Fritts et al. 2006]
MERs [Withers and Smith 2006] Phoenix [Withers and Catling 2010]

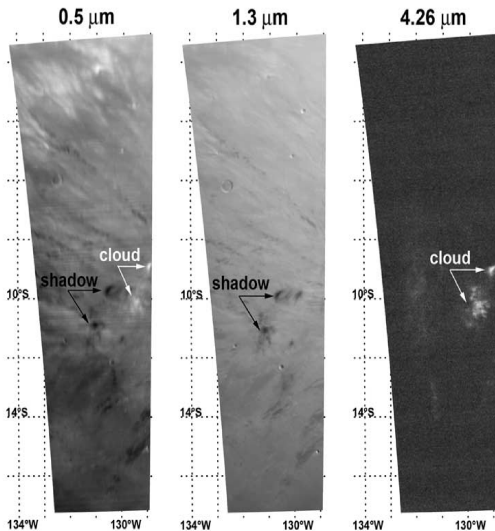


[Schofield, Science 1997; Magalhaes et al. JGR 1999]

Mesospheric CO₂ clouds



[MEx HRSC, Määttä et al. Icarus 2010]

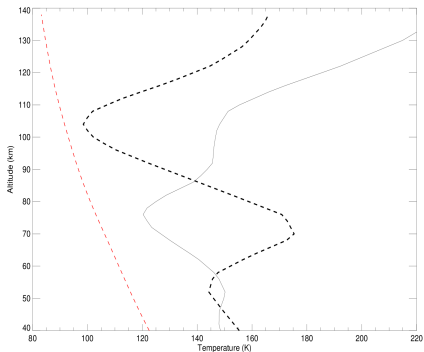


[MEx OMEGA, Montmessin et al. JGR 2007]

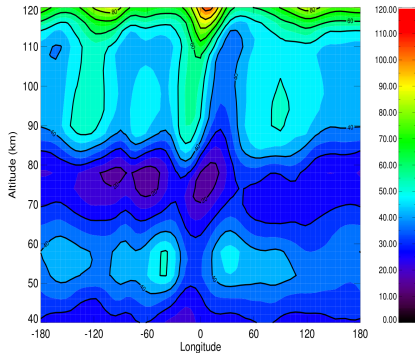
Global Circulation Modeling

Role of thermal tides in the formation of cold pockets propitious to CO₂ clouds

T profiles night (dash) / day (full)



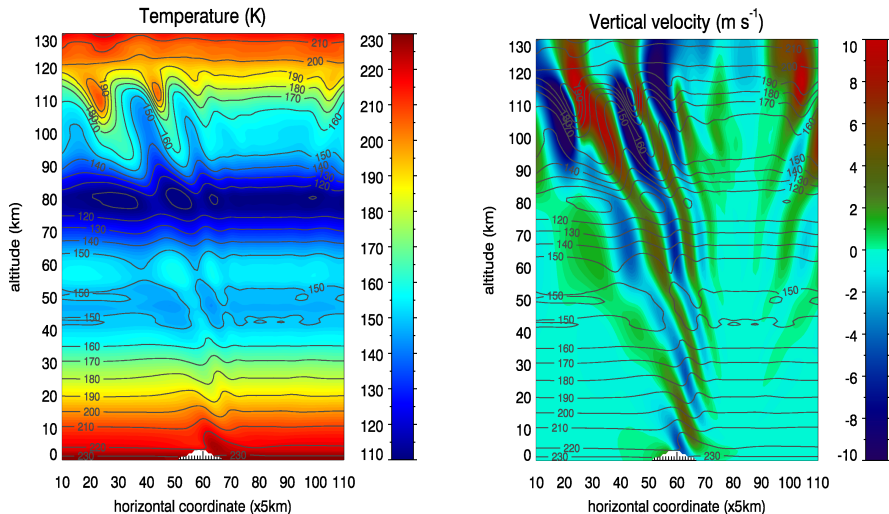
Spatial variations T - T_{sat}



[L_S is 0 – 30°, latitude and longitude 0°. Gonzalez-Galindo et al. Icarus 2011]

3D GW simulation with 30 m s^{-1} rightward wind

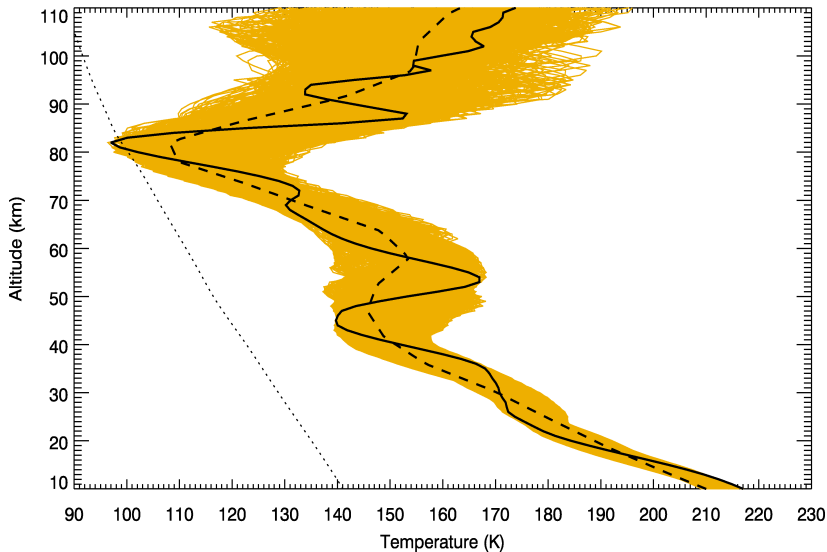
$\delta x = 5 \text{ km}$, $\delta z \sim 1 \text{ km}$, model top 180 km with 50-km sponge layer



[Spiga et al. GRL 2012]

Gravity waves & subcondensation pockets

Full: Large-scale profile. Dashed and envelope: + resolved mesoscale waves

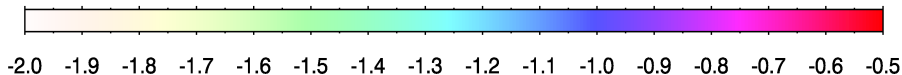
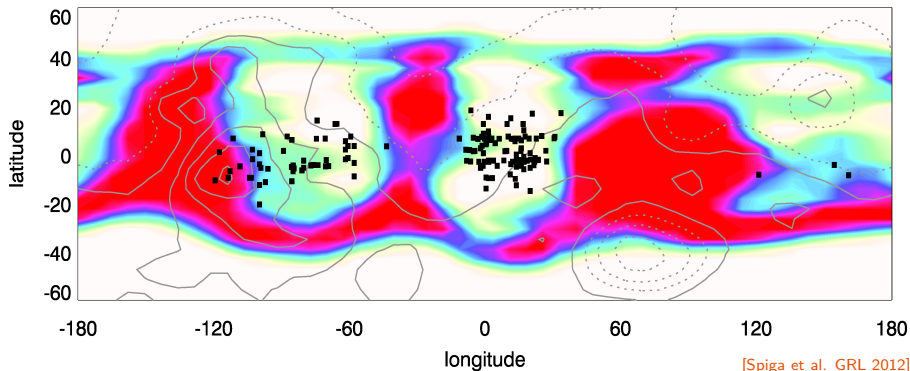


[Spiga et al. GRL 2012]

Spatial variations of GW filtering \rightarrow \mathcal{S} maps

Northern spring CO₂ clouds

Regions/seasons with observed mesospheric CO₂ clouds feature propitious atmospheric conditions for GW propagation.

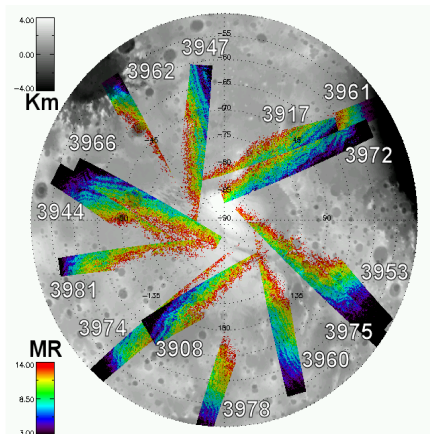


Outline

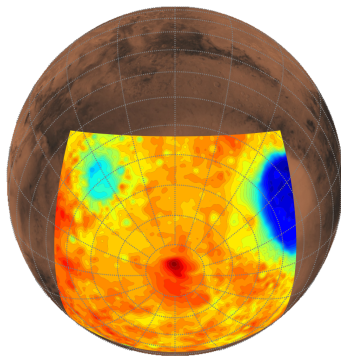
- 1 Methodology
- 2 Slope winds
 - Tharsis volcanoes
 - Polar regions
- 3 Boundary layer convection
 - Vortices
 - Radiative control
- 4 Gravity waves
 - Mesosphere
 - Troposphere
- 5 Rocket dust storms
- 6 Conclusion

Mapping of gravity waves by O₂ airglow

OMEGA airglow observations



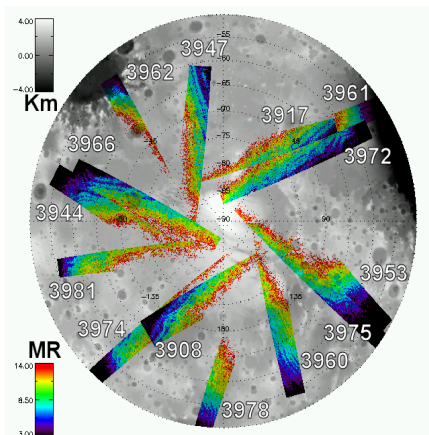
Mesoscale polar modeling



[Altieri et al. JGR 2012]

Mapping of gravity waves by O₂ airglow

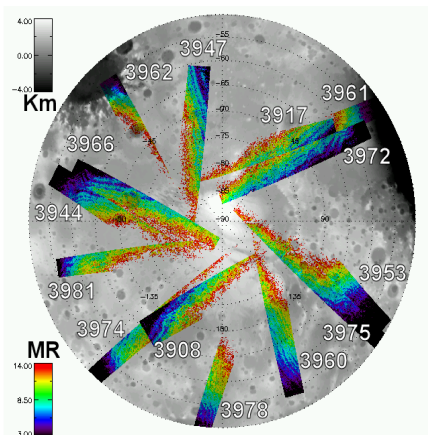
OMEGA airglow observations



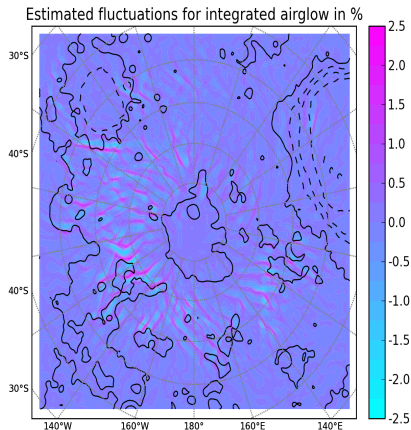
Mesoscale polar modeling

Mapping of gravity waves by O₂ airglow

OMEGA airglow observations



Mesoscale polar modeling



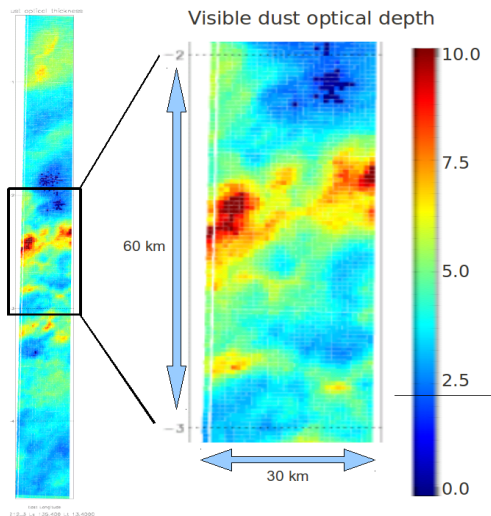
[Altieri et al. JGR 2012]

Outline

- 1 Methodology
- 2 Slope winds
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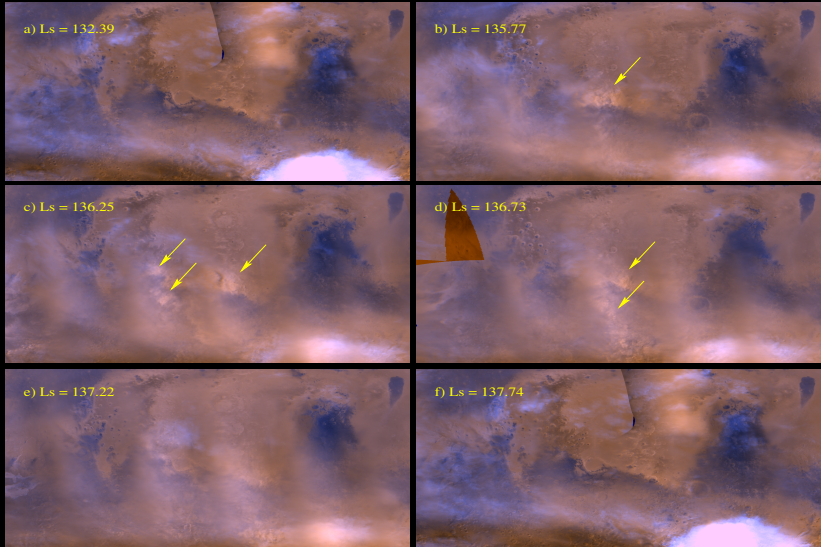
The “OMEGA storm” witnessed by Mars Express

A complex, cumuliform, dust storm in Terra Meridiani at $L_S = 135^\circ$



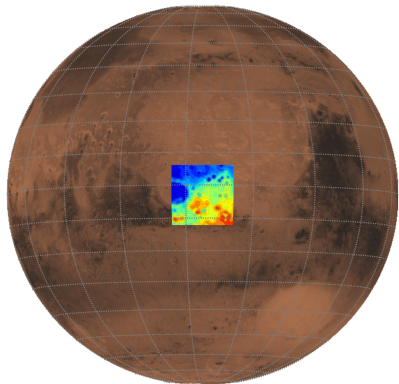
Adapted from Määttä et al. *Icarus* 2009

MOC images before, during, after “OMEGA” storm



[Image built from Mars daily global images downloaded in the “Mars Climate Center” website hosted by Ashima Research]

Mesoscale simulation

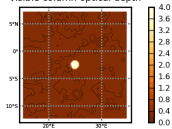


- LMD-MMM with tracers [*Spiga & Forget JGR 2009*]
- Dust radiative transfer and 2-moment transport scheme [*Madeleine et al. JGR 2011*]
- Recent dust optical indices [*Wolff et al. JGR 2009*]

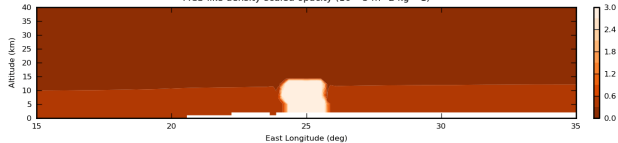
- Terra Meridiani site [OMEGA]
- $181 \times 181 \times 101$ grid points
- 7 km horizontal grid spacing
- ~ 700 m vertical grid spacing with model top at 1 Pa

Afternoon. Local time 1400

Visible column optical depth



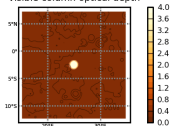
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



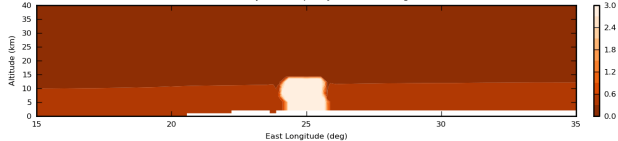
[Spiga et al. JGR 2012, arxiv 1208.5030]

Afternoon. Local times 1400, 1600, 1800

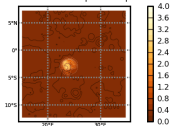
Visible column optical depth



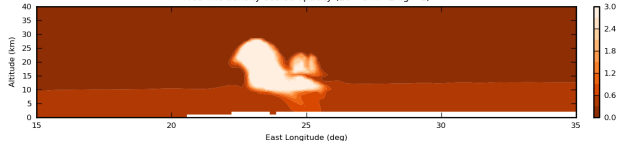
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



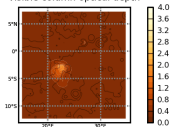
Visible column optical depth



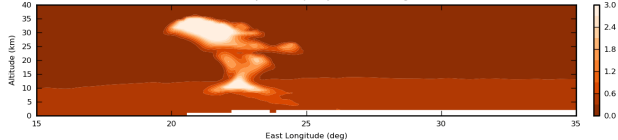
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



Visible column optical depth



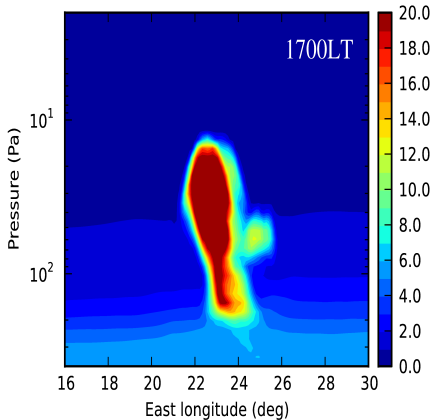
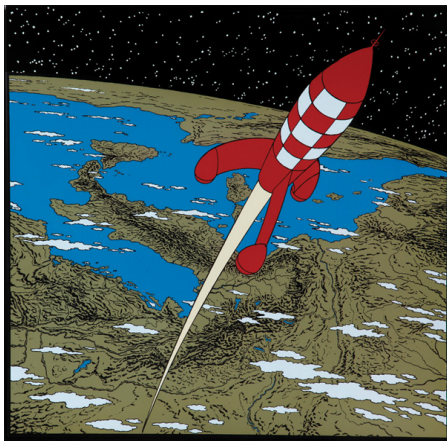
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



[Spiga et al. JGR 2012, arxiv 1208.5030]

“Rocket dust storms”!

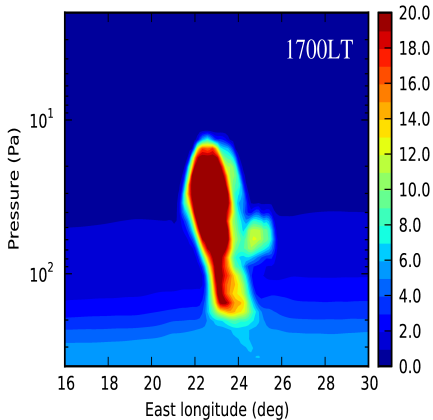
Rapid and powerful vertical transport of dust particles



[Left picture extracted from Hergé Casterman 1954]

... or: “conio-cumulonimbus”

Dust-driven deep convection on Mars

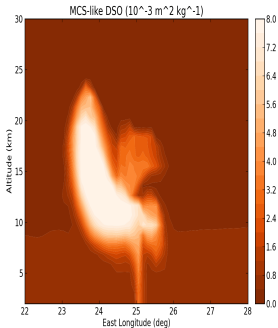


[Left picture downloaded from NOAA website]

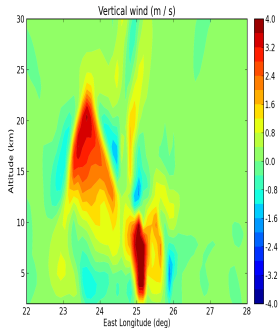
Rocket dust storm [a.k.a. conio-cumulonimbus]

Dust-driven deep convection on Mars

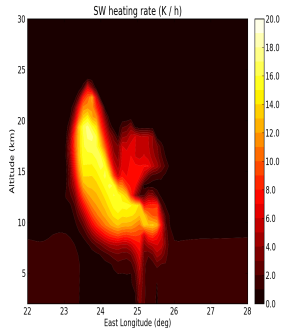
(DS) Optical depth



Vertical wind



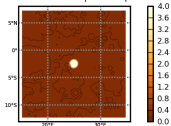
SW heating rate



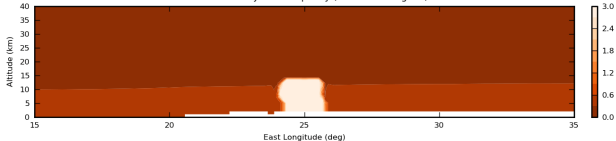
[Spiga et al. JGR 2012, arxiv 1208.5030]

Afternoon. Local times 1400, 1600, 1800

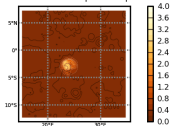
Visible column optical depth



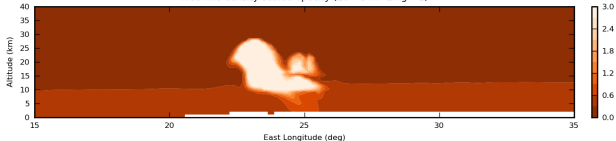
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



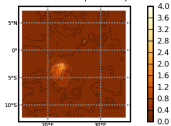
Visible column optical depth



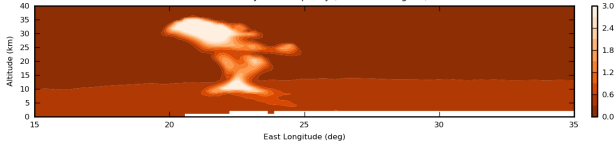
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



Visible column optical depth



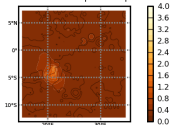
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



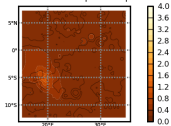
[Spiga et al. JGR 2012, arxiv 1208.5030]

Evening. Local times 2000, 2200, 0000

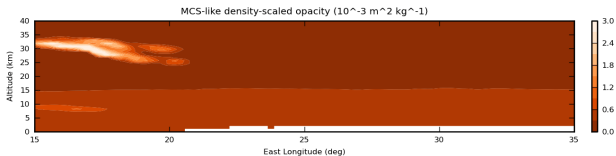
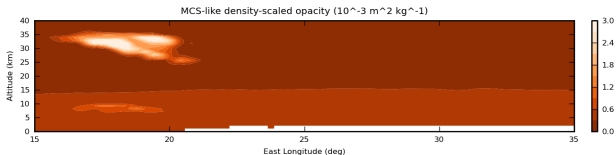
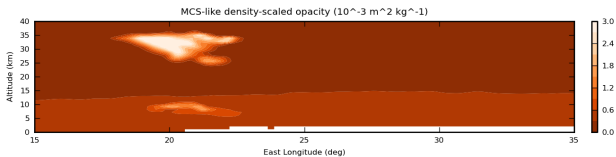
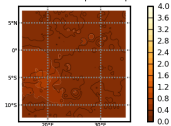
Visible column optical depth



Visible column optical depth

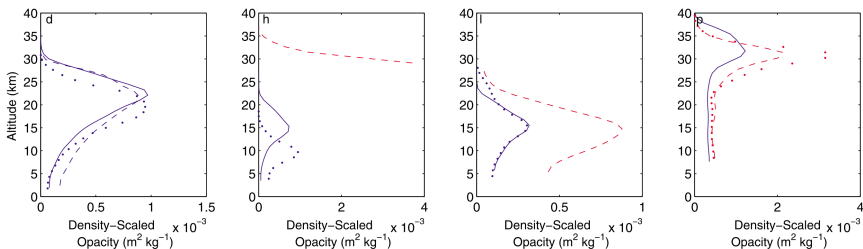
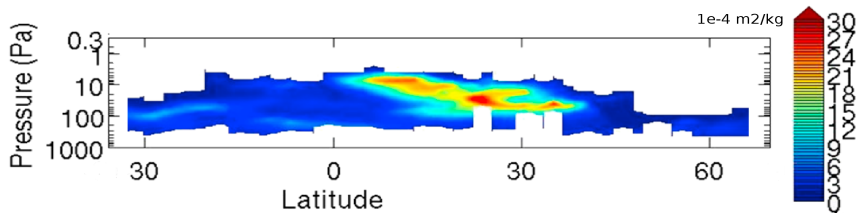


Visible column optical depth



[Spiga et al. JGR 2012, arxiv 1208.5030]

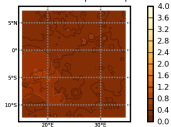
Detached layers of dust: MCS observations



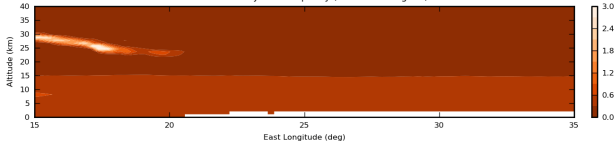
[Heavens et al. JGR 2011 (part 1 & 2)]

Nighttime. Local times 0200, 0400, 0600

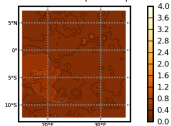
Visible column optical depth



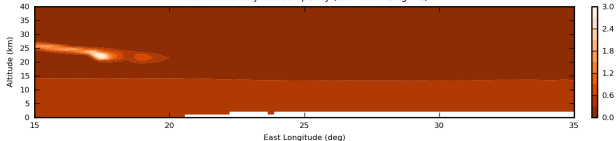
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



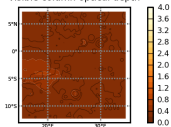
Visible column optical depth



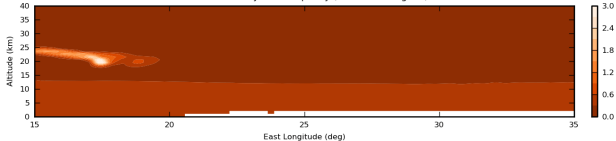
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



Visible column optical depth



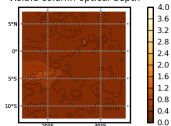
MCS-like density-scaled opacity ($10^{-3} \text{ m}^{-2} \text{ kg}^{-1}$)



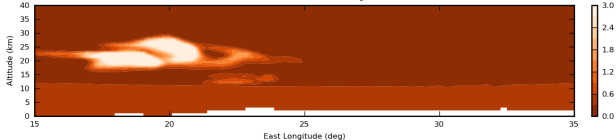
[Spiga et al. JGR 2012, arxiv 1208.5030]

Morning. Local times 0800, 1000, 1200

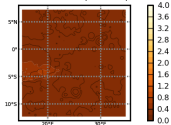
Visible column optical depth



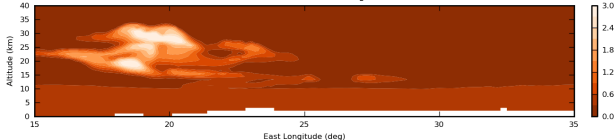
MCS-like DSO ($10^{-3} \text{ m}^2 \text{ kg}^{-1}$)



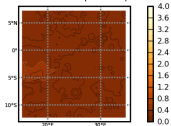
Visible column optical depth



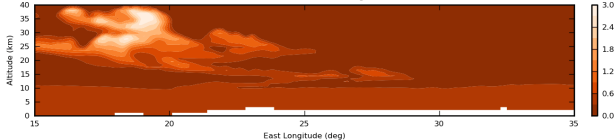
MCS-like DSO ($10^{-3} \text{ m}^2 \text{ kg}^{-1}$)



Visible column optical depth

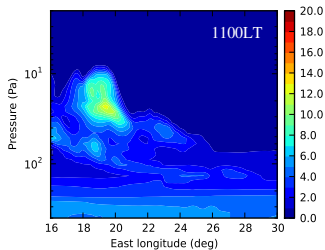
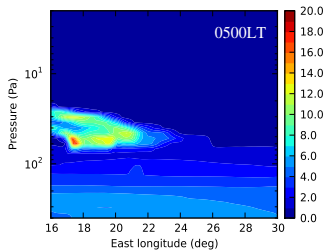
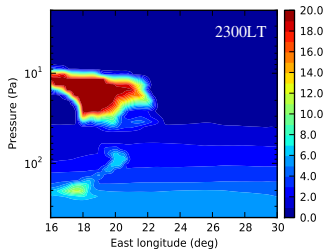
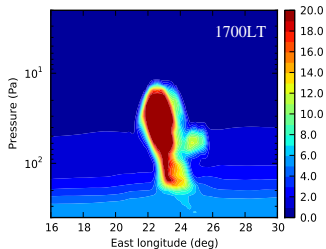


MCS-like DSO ($10^{-3} \text{ m}^2 \text{ kg}^{-1}$)



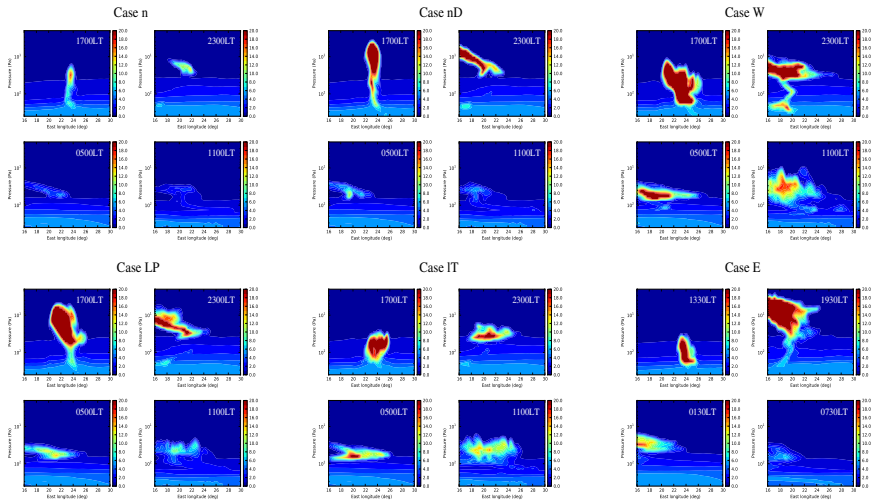
[Spiga et al. JGR 2012, arxiv 1208.5030]

From rocket dust storm to detached dust layers



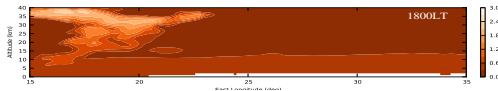
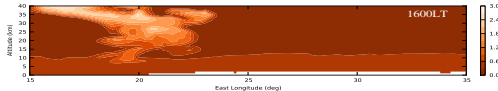
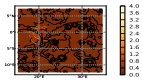
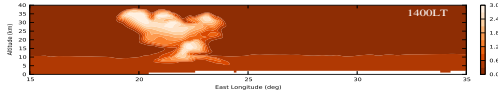
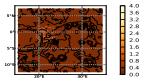
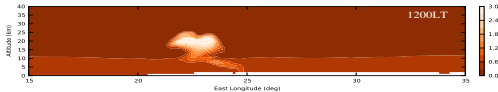
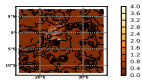
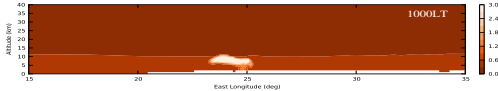
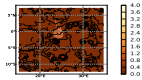
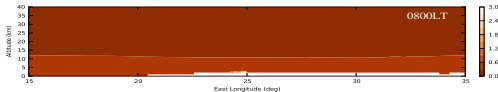
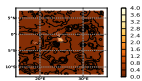
[Spiga et al. JGR 2012, arxiv 1208.5030]

Sensitivity study: results



[Spiga et al. JGR 2012, arxiv 1208.5030]

With simplified lifting (storm area only, $\sigma_t = 5 \text{ mN m}^{-2}$, $\alpha = 2 \times 10^{-3} \text{ m}^{-1}$)



Factor of variability for rocket dust storms

Major

- **Incoming solar radiation:** the further the dust disturbance from subsolar latitudes, the weaker the convection.
- **Background dustiness:** the less dusty the atmosphere, the lower the environmental temperature, the stronger the convection.

Less crucial

- **Atmospheric lapse rate:** A more stable profile would tend to inhibit convective ascent of rocket dust storms.
- **Wind shear:** strong vertical shear impact vertical extent.

Necessary conditions

- Dust particles available for lifting.
- Lifting threshold reached.

Variability of subsequent detached layers of dust

High-altitude detached dust layers by rocket dust storms would

- ☞ be particularly discernible in low-latitudes from late N winter to late N summer (“high-altitude tropical dust maximum”)
- ☞ correlate with interannual variability of “early dust storm season”
- ☞ yield significant longitudinal variability (fast mesoscale processes)
- ☞ create enriched layers within background dust layer in dusty season, instead of “truly detached” layers in clear seasons

⇒ scenario consistent with MCS measurements

[McCleese et al. JGR 2011, Heavens et al. JGR 2011 part 1 & 2]

Implications of dusty deep convection on Mars

In addition to the impact on dust distribution:

- ☞ importance of mesoscale processes
- ☞ impact on global circulations (heat & momentum budget, predictability, planetary waves); GCM parameterization needed!
- ☞ impact on regional/global dust storms and their dynamics;
- ☞ vertical transport of water vapor and chemical species;
- ☞ generation of strong electric fields;
- ☞ source of gravity waves;
- ☞ atmospheric hazard for robotic and human exploration;
- ☞ comparative planetology perspectives.

Outline

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Rich mesoscale and microscale meteorology!

- ➡ Powerful slope winds \Rightarrow surface budget: heat, volatiles, ...
- ➡ PBL convection: not-so-shallow, radiatively controlled (\neq Earth)
- ➡ Gravity waves: troposphere and mesosphere (\Rightarrow CO₂ clouds)
- ➡ Rocket dust storms: radiatively-induced deep convection

Selected references and contact

- ➡ JGR 2009, QJRMS 2010, Icarus 2011, GRL 2012, JGR 2013
- ➡ PDFs available <http://www.lmd.jussieu.fr/~aslmd>
- ➡ E-Mail aymeric.spiga@upmc.fr
- ➡ Twitter @aymeric_spiga